



Investigation of Carbon Dioxide mystery in the atmosphere of Mars using oxygen triple isotope anomaly.

Vera Smirnova¹, Robina Shaheen², Mark Thiemens²

Department of NanoEngineering¹, Department of Chemistry and Biochemistry² University of California San Diego.

Abstract and Background

The red planet - Mars, today is a frigid cold desert (frozen CO₂ and H₂O) with thin air (~6mbar). However, the presence of deep channels and weathered rocks documented by various satellite and rovers on Mars serve as evidence that Mars had once experienced liquid water and dense atmosphere. Phoenix data has shown that the heat generated due to a strong UV radiation can trigger huge dust storms and sublimation of CO₂ and H₂O. The interaction of UV light ($\lambda < 180\text{nm}$) with gaseous CO₂ leads to its dissociation into carbon monoxide (CO) and oxygen. Therefore it is expected to see a significant amount of CO in the atmosphere, however the latest Mars space probe MAVEN (Mars Atmosphere and Volatile Evolution mission) has not detected even a trace of CO. Therefore the question stands of where does carbon monoxide go?

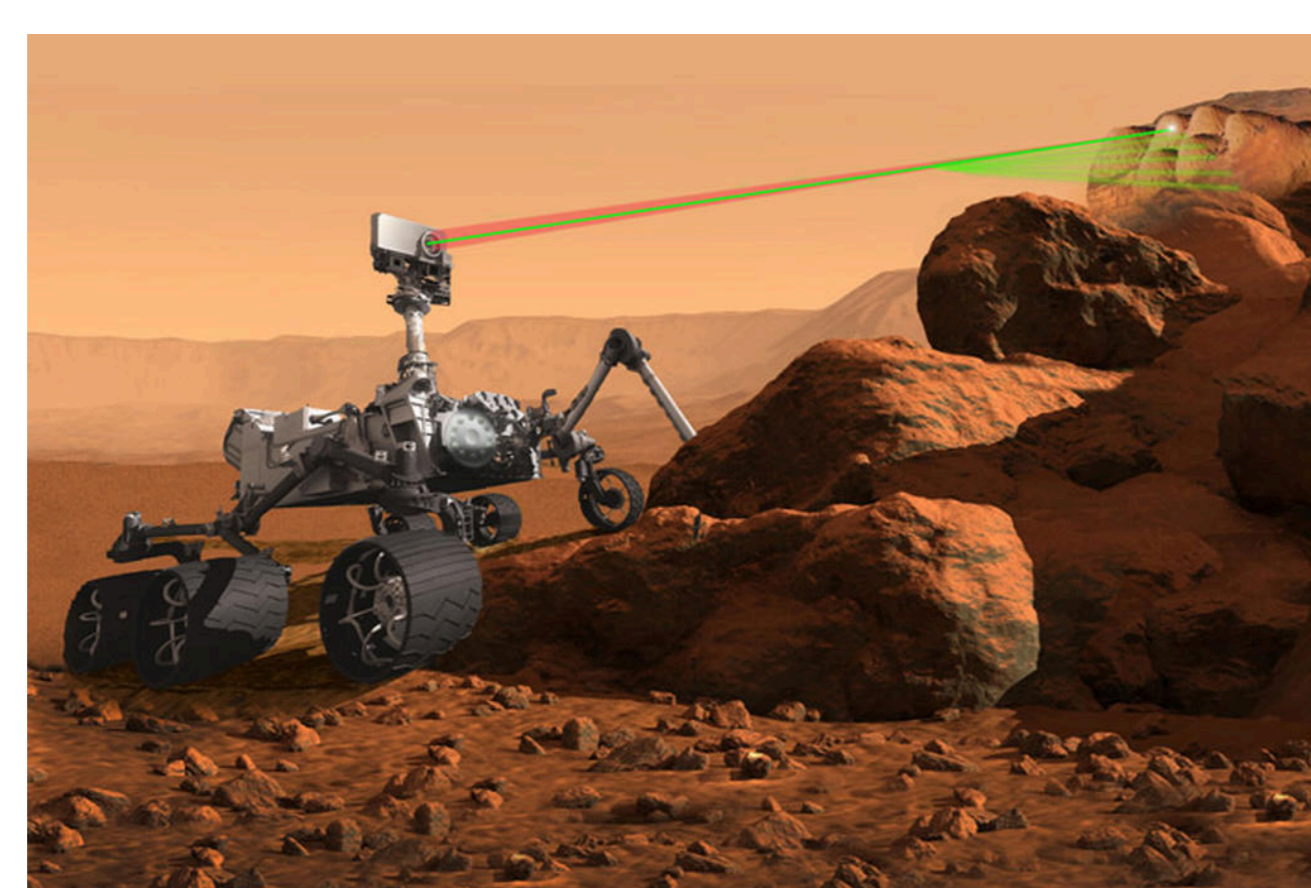
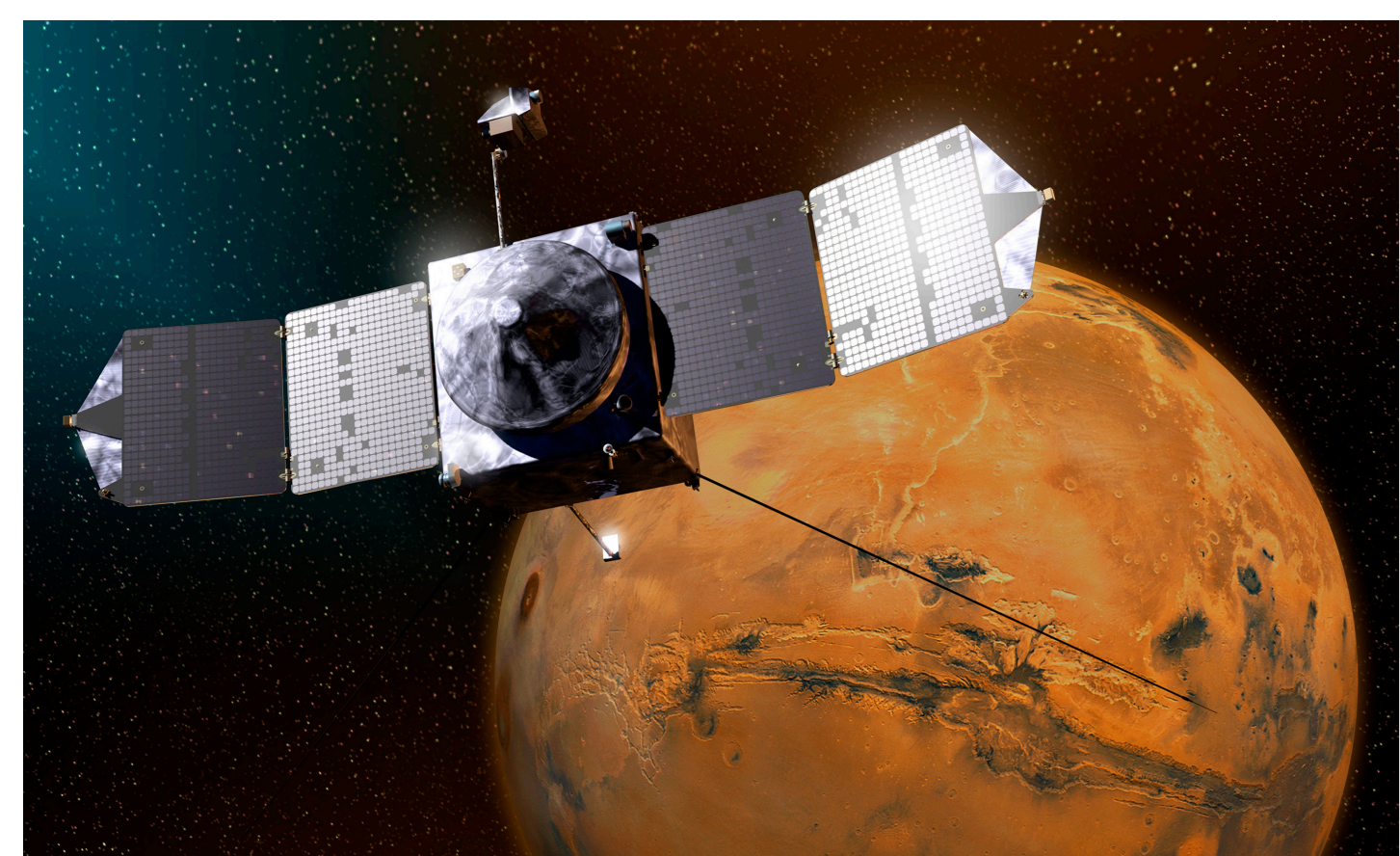


Figure 1a: MAVEN to measure atmospheric composition of Mars

Figure 1b: Curiosity to investigate Martian climate and geology

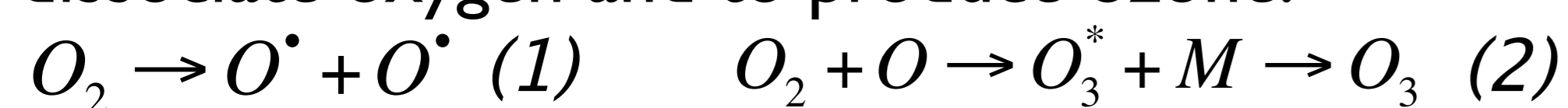
Method

CO and O₂ of known isotopic composition was used to investigate the formation of CO₂ in the presence of UV light. Two sets of experiments were designed to simulate Mars' present day atmosphere. Water was added in Set 1 to produce OH radicals to simulate summer time gas phase photochemistry:

Set 1: CO, O₂ and H₂O

Set 2: CO and O₂

A source of UV light is mercury discharge lamp ($\lambda_1 = 184\text{nm}$, $\lambda_2 = 253\text{nm}$) $\lambda = 184\text{nm}$ dissociate oxygen and to produce ozone:



$\lambda = 253\text{nm}$ dissociates ozone molecules to O(¹D) which may react with CO to form CO₂.

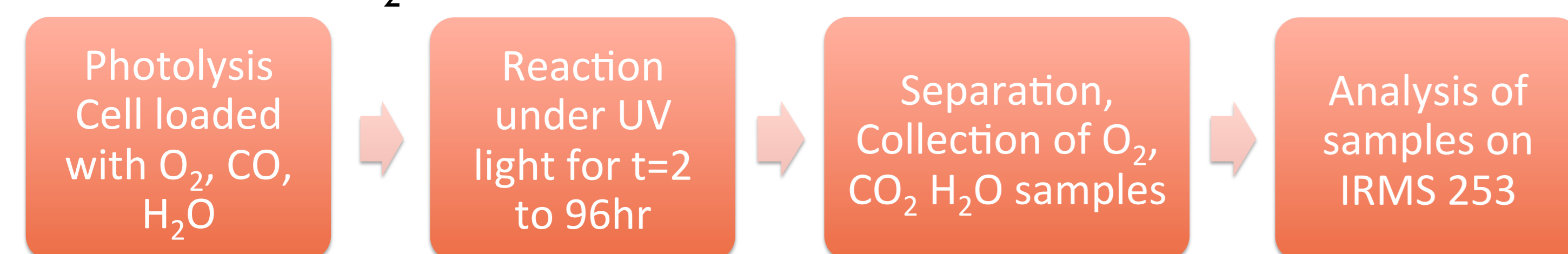


Figure 2: Schematic of the general workflow of the experiment

Many process such diffusion, evaporation follow mass dependent fractionation. The commonly used delta (δ) value denote the isotope ratios ¹⁷O/¹⁶O (¹⁷R) and ¹⁸O/¹⁶O (¹⁸R).

$$\delta^{17}O \text{ permil} = \left(\frac{{}^{17}R_{\text{sample}}}{{}^{16}R_{\text{standard}}} - 1 \right) * 1000$$

$$\Delta^{17}O = \delta^{17}O - 0.52\delta^{18}O$$

Experimental Set Up

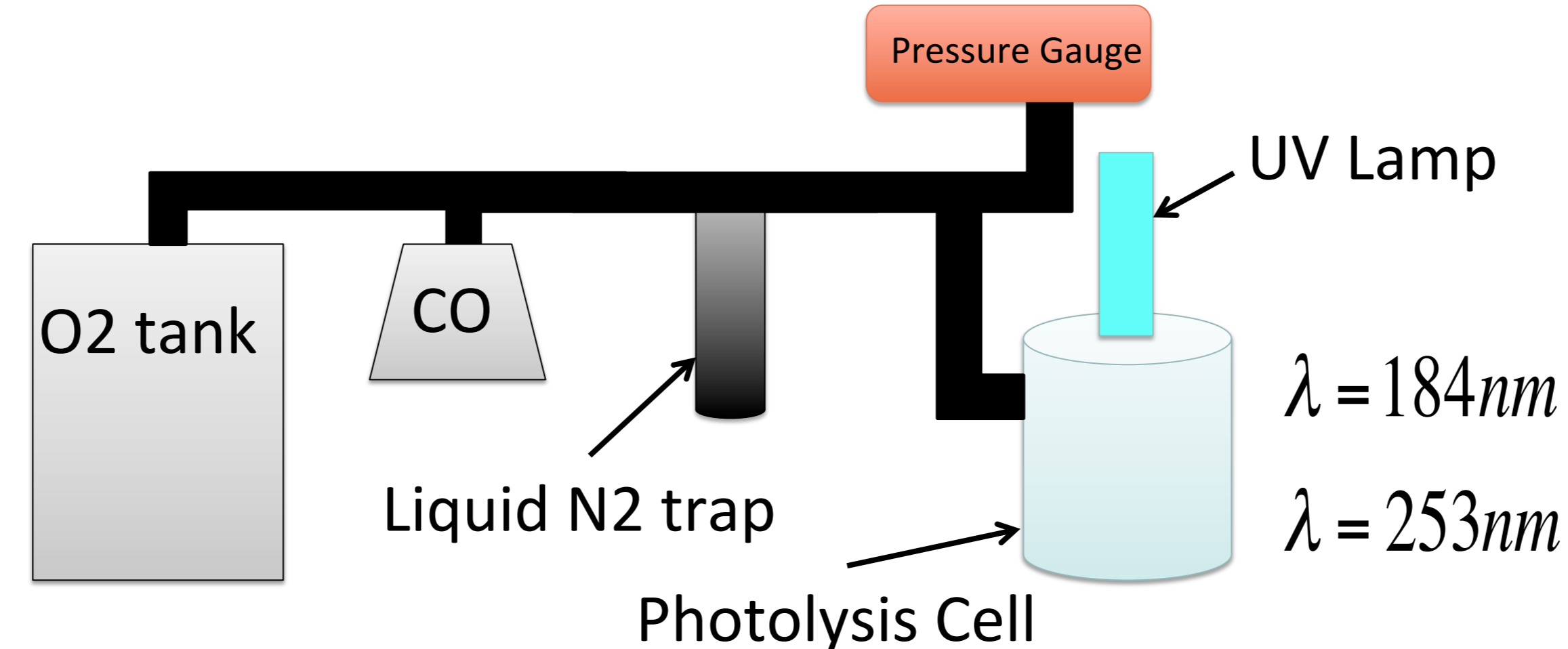


Figure 2: Visual graphic of the experimental set up

Results

The reaction of CO with O₂ in the presence of H₂O (Set 1 experiments) indicated depletion in the oxygen isotopes with time in product CO₂ and O₂ (Fig. 4a). In Set 2 CO₂ got lighter, no significant change in O₂. A comparison between Set 1 and Set 2 indicated that O₂ exchange isotopically with H₂O through hydroxyl radicals (eq. 4-8). The trend suggests presence of enriched intermediate species (O₂H and H₂O₂). In the presence of UV light peroxide dissociate quickly and therefore the tiny amount present at steady state cannot be analyzed.

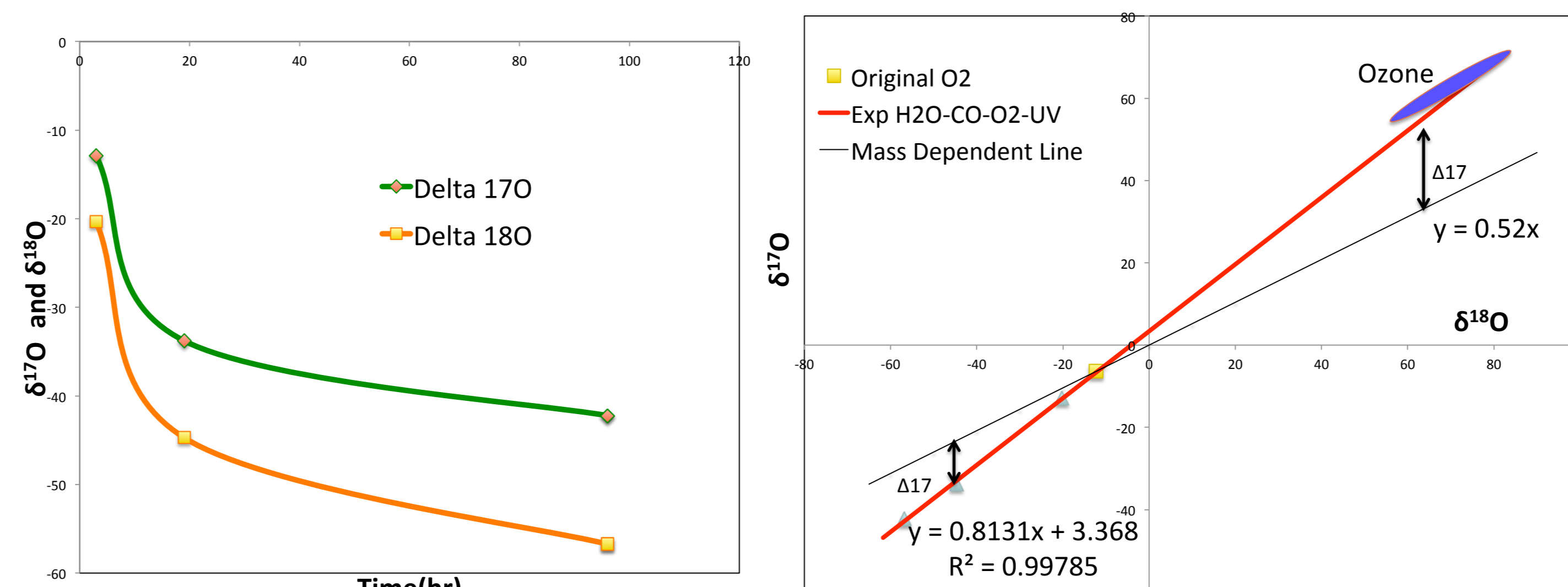
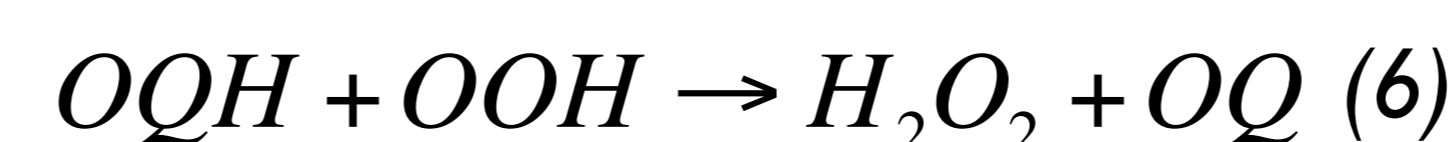
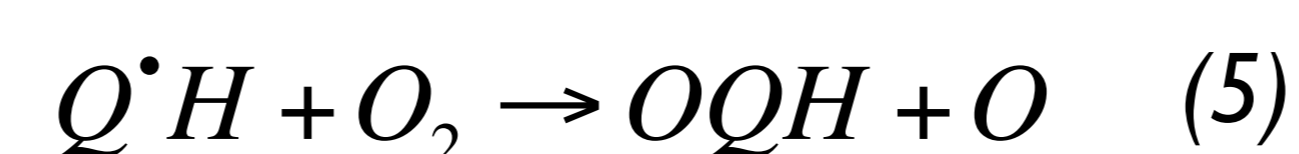
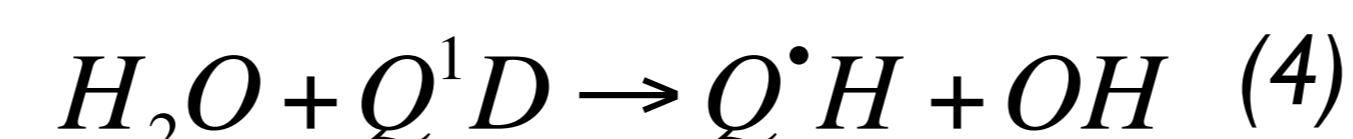


Figure 4a: Graph of $\delta^{17}O$ and $\delta^{18}O$ change with time

Figure 4b: Graph of O₂ fractionation during H₂O-CO-O₂-UV Exp.

Currently we know that the $CO_2 \rightarrow CO + O$ (3) reaction occurs on Mars when UV light of wavelength less than 180 nm hits abundant CO₂ particles in the atmosphere however as there is no build up of CO detected, it has to quickly transform back into CO₂.

Mechanism 1



There are two ways CO can then proceed to CO₂



Mechanism 2:



Discussion

Most of the minerals in meteorites from Mars showed oxygen isotope anomaly in $\Delta^{17}O_{\text{carbonates}} = 0.6-0.9\text{‰}$ and $\Delta^{17}O_{\text{silicates}} = 0.2-0.4\text{‰}$ (Farquhar, Thiemens 2000). Here I have shown that CO₂ exchange isotopically with oxygen and hydroxyl radicals and this isotopic signature can be preserved in the carbonate product from anomalous CO₂ and H₂O. The oxygen isotope anomaly is derived from ozone. The ozone molecule is a natural trace with highest enrichment and large oxygen isotope anomaly (30-40‰) which depends on temperature and pressure (Thiemens, Heidenreich 1983). Large isotope anomalies has been observed in other oxygen carrying molecules in the atmosphere of earth such as CO₂, SO₄, NO₃ and H₂O₂ (Thiemens, Shaheen 2014). For Mechanism 2 the experiments have already been completed, however the data is still being worked on to fluorinate the samples of CO₂ to find the Delta 17O values. Data collected so far is presented in the table below.

Experiment #	Time(hr)	Ratio of O ₂ to CO ₂	$\delta^{13}C$	$\delta^{18}O$
Initial	0	0	-43.00	51.00
1	2	19.38	-36.485	20.996
2	2	29.08	-34.542	19.066
3	2	63.33	-33.654	13.933

Table 1: Changes in the oxygen isotopic composition of CO₂ with ratio of O₂ to CO₂

In my experiment ozone was formed due to presence of UV light and by looking what isotope numbers are of resulting species I can say that, if the original CO was reacting with O(¹D) from ozone it should move up the line, however as shown in the table 1, $\delta^{18}O$ isotope numbers in CO₂ are decreasing which predicts that CO₂ was formed by the reaction with O(³P) instead.

References

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Acknowledgments

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